Vegetation history and monsoonal fluctuations during the last 12,500 years BP inferred from pollen record at Lower Subansiri Basin, Assam, Northeast India

S.K. BERA^{*} AND S.K. BASUMATARY

Birbal Sahni Institute of Palaeobotany, 53 University Road, Lucknow 226 007, India. *Corresponding author: skbera 2000@yahoo.com

(Received 15 July, 2011; revised version accepted 7 November, 2012)

ABSTRACT

Bera SK & Basumatary SK 2013. Vegetation history and monsoonal fluctuations during the last 12,500 years BP inferred from pollen record at Lower Subansiri Basin, Assam, Northeast India. The Palaeobotanist 62(1): 1–10.

A 120 cm sediment profile from 4.5 m deep exposed section of Lower Subansiri River Basin has been pollen analyzed to trace palaeovegetation and climate in relation to monsoonal activity since 12,500 years BP. The study has depicted the existence of tropical mixed tree–savannah type to tropical mixed deciduous type forest under four climatic regimes, viz. cold and dry to warm and relatively dry since 12,500 years BP followed by a palynologically barren zone indicating strong fluvial activity. Four sets of monsoon fluctuations, i.e. relatively low to considerably decline in monsoonal activities have been estimated by using pollen marker taxa during the period. The marker taxa signifying high monsoonal activity are – Dipterocarpaceae, *Syzygium cumunii*, Arecaceae and Moraceae. *Melastoma malabathricum*, the only taxa which is marked as low monsoonal activity and dry depositional environment. The consistent occurrences of high land taxa are suggestive of long distance pollen transport from higher elevation via strong wind and water source.

Key-words-Lower Subansiri Basin, Monsoonal activity, Northeast India, Vegetation.

निचली सुबांसिरी द्रोणी, असम, पूर्वोत्तर भारत के पराग अभिलेख से अनुमानित पिछले 12,500 वर्षों पूर्व के दौरान वनस्पति इतिहास एवं मानसूनी उतार-चढ़ाव

एस.के. बेरा एवं एस.के. बसुमतारी

सारांश

निचली सुबांसिरी नदी द्रोणी के 4.5 मीटर गहरे खंड से अनावरित एक 120 से.मी. अवसाद परिच्छेदिका 12,500 वर्षों पूर्व से मानसूनी गतिविधियों के संबंध में पुरावनस्पति व जलवायु का पता लगाने को पराग विश्लेषित की गई है। अध्ययन ने प्रबल नदीय गतिविधि इंगित करते हुए परागाणविक रूप से अनुत्पादक मंडल के अनुगामी 12,500 वर्षों पूर्व से चार जलवायवी प्रवृत्तियों अर्थात शीत व शुष्क से कोष्ण एवं सापेक्षतया शुष्क के अंतर्गत उष्णकटिबंधीय मिश्रित वृक्ष-सवन्ना प्रकार से उष्णकटिबंधीय मिश्रित पतझड़ी प्रकार की मौजूदगी चित्रित की है। इस अवधि में पराग चिहनक टैक्सा प्रयुक्त करते हुए मानसून उतार-चढ़ावों के 04 समुच्चय अर्थात निम्न के सापेक्षतया पर्यात्त रूप से घटती मानसूनी गतिविधियां अनुमानित की गई हैं। चिहनक टैक्सा डिप्टेरोकार्पेसी, *सायज़ीजियम कम्युनिआई*, एरेकेसी व मोरेसी उच्च मानसूनी गतिविधियां व्यक्त कर रही है। मेलास्टोमा मालाबाधीकम, एक मात्र टैक्सा है जो निम्न मानसूनी गतिविधि एवं शुष्क निक्षेपणीय पर्यावरण के रूप में चिहनित है। उच्च भूमि टैक्सा की निर्बाध प्राप्ति तेज हवा व जल स्रोत के माध्यम से उच्च ऊंचाई से लंबी दूरी पराग परिवहन को सुझावित है।

© Birbal Sahni Institute of Palaeobotany, India

INTRODUCTION

etailed and systematic study of palynological record has provided valuable insights concerning the changing vegetation scenario and contemporaneous climatic changes in the Indian subcontinent during the Quaternary Period (Meher-Homji & Gupta, 1999). Hitherto, no adequate palynological information is available from northeast India. However, some preliminary works have been carried out in Assam (Gupta, 1971; Bhattacharya & Chanda, 1988, 1992; Bera, 2000, 2003; Bera et al., 2008), Manipur (Roy & Chanda, 1987; Nautiyal & Chauhan, 2009), Mizoram (Chauhan & Mandaokar, 2006; Mandaokar et al., 2008), Tripura (Goswami, 1981; Prasad, 1986, 1988; Prasad & Ramesh, 1983), Arunachal Pradesh (Bhattacharyya et al., 2007), Meghalaya (Gupta & Sharma, 1985; Basumatary & Bera, 2007, 2010). Nagaland (Bera et al., 2010) and Sikkim (Bhattacharya & Chanda, 1986; D' Costa & Mukharjee, 1986; Sharma & Chauhan, 2001; Chauhan & Sharma, 1996; Mehrotra et al., 2005).

However, the estimation of monsoonal activity by palynodata has been carried out by few workers in India (Swain *et al.*, 1983; Prasad & Enzel, 2006; Demske *et al.*, 2009). Study of the Indian monsoonal activity is critical for understanding regional and global variability as it drives major changes of the tropical Indian Ocean region (Gupta *et al.*, 2003; Ely *et al.*, 1993; Fleitmann *et al.*, 2003; Goodbred & Kuehl, 2000). The changes in the monsoonal activity over South Asia affect the fluvial system and thus flora and fauna of the region (Gupta *et al.*, 2006; Kumaran *et al.*, 2005). Annual precipitation is an important parameter explaining the modern distribution of pollen taxa is clearly evidenced by distinct pollen assemblage, e.g. *Glochidion* pollen type, *Grewia, Schleichera oleosa, Trema orientalis, Lannea* type, *Maytenus* type, *Lagerstroemia* and Melastomaceae/Combretaceae are receiving annual rainfall 1200 to 2500 mm/yr, *Syzygium* type, *Trema orientalis, Olea dioica, Elaeocarpus, Securinega* type, *Tetrameles nudiflora, Mallotus* type, Moraceae/Urticaceae receiving 2800 to 4500 mm/yr and *Poeciloneuron* and *Syzygium* type are receiving exceeding 4500 mm/yr (Barboni & Bonnefille, 2001).

The present area of study located in Dhemaji District, situated in the northeastern corner of Assam, lies between lat. 27°27' and 27°57' N and long. 94°18' and 95°32' E respectively (Fig. 1). It is bounded in the north by Siang and Papumpare District of Arunachal Pradesh, in the east by the Arunachal Hills, in the west by Subansiri River, North Lakhimpur District and in the south by Brahmaputra River. The Subansiri River (Fig. 2a) is the largest tributary of the mighty Brahmaputra, a Trans-Himalayan river originating from the western part of Mount Pororu (5059 masl) in the Tibetan Himalaya (Goswami et al., 1990). The different geo-ecological zones transversed by the river have distinctive assemblage of topographical, geological, climatological and floral characteristics. After flowing for around 190 km through Tibet and 200 km through Arunachal Himalaya, it enters into the plains of Assam through a gorge near Gerukamukh which is situated in Dhemaji District and flows through North Lakhimpur District and then merges with mighty Brahmaputra River. The riverbed is composed of sand mixed with pebbles and boulders up to Chauldhoaghat (Goswami, 1997). The forest vegetation is tropical mixed deciduous type and the majority of the forest embraces semi-evergreen taxa and mixed deciduous taxa including riparian and swamp taxa (Champion & Seth, 1968). Floristically, due to the excessive deforestation, the forest areas are cleared and secondary monoculture forests have gradually been established. These forests mainly consist of Shorea assamica, Mesua ferrea, Cinnamomum



Fig. 1—Map showing the study area.

BERA & BASUMATARY-VEGETATION HISTORY AND MONSOONAL FLUCTUATIONS AT LOWER SUBANSIRI BASIN, ASSAM 3



Fig. 2-(a) A view of Lower Subansiri River Basin, near NHPC Ltd., Assam. (b) A view of vegetation composition, Lower Subansiri River Basin, Assam.

bejolghota, Dipterocarpus macrocarpus, Dillenia pentagyna, Lagerstroemia parviflora, Schima wallichii, Aporusa wallichii, Albizia lebbeck, Terminalia myriocarpa and Tetrameles nudiflora (Fig. 2b). The main components of shrubby species are Melastoma malabathricum, Holarrhena antidysenterica,

Garcinia lancifolia and Acacia concinna. Palms like Areca, Caryota and Pinanga are also conspicuous. Bamboos such as Bambusa arundinaria, Dendrocalamus hamilitonii and Melocanna bambusoides are common in and around of the study area. The ground flora in deciduous forest is very poor



LEGENDS-

- A₁- Newer alluvium; A₂-Older alluvium;
- B-Boulder, pebble and sand with little silty clay (Fluvial zone);
- C-Fine sand with little pebble; D-Fine sand;
- E-Fine sand with pebble and wood fossil; ●-Sample for C¹⁴ date; □ -Wood fossil

Fig. 3-Lithostratigraphy shown in photograph (left) and sketches (right) of Subansiri river section, Assam.

and seasonal. Ferns and fern-allies, liverworts and mosses are also seen on old tree trunks, rock, boulders, etc. along the water streams and shady places. The natural reserve forest is highly exposed to large scale deforestation primarily due to fulfil the needs of timber, firewood and ongoing construction of 2000 mw Lower Subansiri Dam by National Hydroelectric Power Corporation (NHPC Ltd). However, the over exploitation of natural forests are resulting in spreading of secondary degrade forest types which support inferior quality of trees and alien weeds.

The climate is generally hot and sultry during summer and moderate during winter. In general, temperature rises maximum upto 37° C during summers and minimum upto 2° C during winters. The relative humidity reaches high up to 90% during rainy season (Forest working plane, 2005, Dhemaji District, Assam). During the months of May to September, there is an average annual rainfall of 3,435 mm. Under the influence of south west monsoons, huge volumes of floodwater start spilling all along the 720 km length of embankments of the River Brahmaputra in Assam, India (Goswami, 1997).

The soil of the region is broadly classified into four groups, namely new alluvial, old alluvial, red loamy and lateritic soil. The new alluvial soil is found in the flood areas subjected to occasional flood and consequently receives annual silt deposit when the flood recedes. Geologically, older and newer alluvium soil covers the area. Piedmont deposits of older alluvium consist of boulders, cobbles, gravel, sand and silt. These are the main repository in the foothill zone. Flood plain area with newer alluvium consists of fluvially derived sediments including gravel, pebble, coarse to medium sand, silt and clay received from the rivers coming from the upper reaches are the main deposits next to the piedmont deposits. The old alluvial soils are developed at higher level and are not subjected to flooding. Red loamy soils have formed on hill slopes under high rainfall conditions. However, soil in the interior forest reserve is deep red clay with alkaline nature (Forest working plan, 2002, Dhemaji District, Assam). The detailed lithostratigraphic succession of the Subansiri River Section has been given (Fig. 3). The upper portion of lithocolumn (A, A, and B) belongs to the older alluvium and rest part (C, Dand E) is under the Tipam Group (Middle-Late Miocene) as detected the scattered silicified fossil woods (Karunakaran, 1974). This study communicated the age-constrained pollen record from Lower Subansiri River Basin which is for the first time to trace palaeovegetation and palaeomonsoon activity along with palaeoflood in and around the study area.

MATERIAL AND METHOD

A 4.5 m deep exposed Subansiri River Section was examined for interpretation of palaeovegetation, climate and monsoonal activity. The upper 1.7 m exposed section has been sampled at 10 cm intervals for palynological studies. However, below 2.8 m depth a Tertiary zone (Tipam Group, Middle-Late Miocene in age) was detected as evidenced by the scattered silicified fossil woods (Mehrotra et al., 2011). The collected soil samples were chemically processed using standard acetolysis method (Erdtman, 1953), 175-310 pollen grains per sample were identified and later counted. The four samples collected from the depths of 120, 100, 70 and 40 cm (base to top) of sedimentary soil profiles were carbon dated at Birbal Sahni Institute of Palaeobotany, Lucknow, India. The percentage of pollen taxa has been calculated with respect to the total pollen count. The pollen diagram is grouped on the basis of various characteristic plant taxa, viz. arboreal taxa (tree and shrub), nonarboreal taxa (terrestrial, marshy and aquatics) and high land taxa (conifers and other broad leaved taxa) respectively in the palynoassemblage (Pl. 1).

RESULTS

Four pollen zones have been recognized in this pollen diagram based on the pollen assemblage recovered from the sedimentary soil profile. Each pollen zone is prefixed by S

PLATE 1 Palynoassemblage recovered from sedimentary profile from Lower Subansiri River Basin, Assam.

	, j	in the second	,,
1.	Dipterocarpaceae	16.	Tubuliflorae
2.	Lagerstroemia sp.	17.	Artemisia sp.
3.	Syzygium cumunii	18.	Chenopodiaceae
4.	Acacia catechu	19.	Justicia simplex
5.	Moraceae	20.	Cerealia
6.	Dillenia pentagyna	21.	Poaceae
7.	Salmalia malabaricum	22.	Cyperaceae
8.	Arecaceae	23.	Pinus sp.
9.	Sapotaceae	24.	<i>Betula</i> sp.
10.	Mesua ferrea	25.	Alnus sp.
11.	Terminalia bellerica	26.	<i>Nymphaea</i> sp.
12.	Strobilenthes sp.	27.	<i>Typha</i> sp.
13.	Melastoma malabathricum	28.	Lycopodium sp.
14.	Polygonum serrulatum	29.	Davallia sp.

15. Impatiens sp.

- Davallia sp. 29
- 30 Microthyriacae



BERA & BASUMATARY—VEGETATION HISTORY AND MONSOONAL FLUCTUATIONS AT LOWER SUBANSIRI BASIN, ASSAM 5

PLATE 1

after the name of the locality Subansiri from where the soil profile has been pollen analyzed (Faegri & Iversen, 1964). These pollen zones are described separately below (Fig. 4).

Pollen zone S-I (120-100 cm)-Grasses-Cyperaceae-Tubuliflorae-Solanaceae-Nymphaea-Salmalia-Sapotaceae assemblage: This pollen zone is extending from a depth of $120 \text{ cm} (12,500\pm170 \text{ years BP})$ to $100 \text{ cm} (7,960\pm170 \text{ years})$ BP) has a time span of 4,540 years. The lithocolumn is characterized by silty clay with little sand and pebble. This pollen zone is characterized by the predominance of nonarboreal taxa (76.5%) over arboreal taxa (10.8%) and the high land taxa (12.7%) respectively. Among nonarboreal taxa, grasses (wild and cerealia type) are dominant (21.4% and 8.3%), Tubuliflorae (7.2%), Liguliflorae (5.6%), Malvaceae (3.0%), Artemisia (2.9%), Lamiaceae (2.1%), Euphorbiaceae (1.9%), Convolvulaceae (1.6%), Acanthaceae (1.1%), Impatiens, Chenopodiaceae/Amaranthaceae, Apiaceae, Caryophyllaceae and Brassicaceae are recorded as other prominent terrestrial nonarboreals at lower values. However, marshy taxa like Cyperaceae (8.8%), *Polygonum* (3.4%) and Onagraceae (3.2%) are encountered at moderate value. The prominent aquatic taxa like Nymphoides (1.3%), Nymphaea (1.1%), Potamogeton and Typha are recorded scarcely. Combretaceae (1.9%) associated with Mesua ferrea (1.3%) and Schima (1.1%) are represented as major tree taxa, however, Arecaceae, Adina,

Salmalia malabarica, Ilex, Dipterocarpaceae, Moraceae and *Emblica officinalis* are sporadic. The shrubby elements such as Oleaceae, *Melastoma malabathricum* and Fabaceae are also encountered scarcely. The extra–regional taxa like *Pinus* (4.3%), *Betula* (2.9%), *Alnus* (2.6%) and *Quercus* (1.1%) are registered consistently in moderate value. However, *Corylus*, *Ulmus* and Ericaceae are sporadic.

Pollen zone S-II (100-70 cm)-Grasses-Bambusoideae-Caryophyllaceae-Impatiens-Onagraceae-Nymphaea-Potamogeton-Dipterocarpaceae-Terminalia-Lagerstroemia-Ilex-Careya assemblage: This pollen zone extending from a depth of 100 cm $(7,960 \pm 110 \text{ years BP})$ to 70 cm $(6,421 \pm 90 \text{ m})$ years BP) has a time span of 1,539 years. The lithocolumn is characterized by silty clay with little sand. This pollen zone is characterized by the dominance of nonarboreals (61.3%) over arboreals (24.1%) and the high land taxa (14.6%) respectively. Among nonarboreals, grasses (wild and cerealia type) are dominant (11.4% and 8.4%). Tubuliflorae (5.1%), Liguliflorae (4.0%), Artemisia (2.7%), Malvaceae (2.0%), Euphorbiaceae (1.1%), Lamiaceae, Impatiens and Convolvulaceae (1.0% each), Acanthaceae (0.8%), Chenopodiaceae/ Amaranthaceae, Caryophyllaceae and Brassicaceae (0.7% each) and Apiaceae (0.4%) are recorded as other terrestrial nonarboreals at good to moderate values. Similarly, marshy taxa like Cyperaceae (6.8%), Onagraceae (3.8%) and Poly-



LEGENDS- 0-30- Silty clay with rootlets; 30-40- Clay; 40-70- Silty clay; 70-100- Silty clay with little sand; 100-120- Silty clay with little sand and pebble; 120-170- Boulder, pebble and sand with little silty clay



gonum (3.1%) are recorded at good values. Among the aquatic taxa, Nymphoides (2.4%), Nymphaea and Typha (1.5% each) and Potamogeton (1.2%) are moderately recorded. Among major trees, Combretaceae is represented as major constituent (2.4%) followed by Arecaceae (2.0%), Dipterocarpaceae (1.9%), Emblica officinalis (1.7%), Moraceae and Salmalia malabaricum (1.6% each), Schima (1.5%), Ilex (1.2%) and Adina (1.0%) respectively. However, Mesua ferrea, Dillenia, Sapotaceae, Lagerstroemia, Syzygium, Meliaceae and Semecarpus are sporadic. Fabaceae (1.9%) and Oleaceae (1.1%) are moderately recorded. Melastoma malabathricum and Strobilenthes are exhibited scarcely. The high land taxa like Pinus (4.6%), Betula (3.1%), Alnus (2.8%), Corylus (1.2%), Quercus and Ericaceae (1.1% each) are recorded moderately. Ulmus is sporadic.

Pollen zone S–III (70–40 cm)—Grasses–Tubuliflorae– Artemisia–Potamogeton–Dipterocarpaceae–Terminalia– Melastoma assemblage: This pollen zone extending from a depth of 70 cm (6,421 \pm 90 years BP) to 40 cm (4,270 \pm 70 years BP) has a time span of 2,151 years. The lithocolumn is characterized by 30 cm silty clay. This pollen zone is characterized by the predominance of nonarboreals (49.7%) over arboreals (33.7%) along with high land taxa (16.6%). Among nonarboreal taxa, grasses (wild and cerealia type) are dominated at the value of 6.2% and 9.5% respectively. Other terrestrial herbs like Tubuliflorae (3.0%), Liguliflorae (2.9%), Artemisia (1.9%), Impatiens (1.8%) and Malvaceae (1.2%) are recorded moderately. However, Caryophyllaceae, Convolvulaceae, Euphorbiaceae, Chenopodiaceae/Amaranthaceae, Brassicaceae, Apiaceae, Lamiaceae and Acanthaceae are recorded in sporadic frequency. Amongst, the marshy taxa, Cyperaceae (4.7%), Onagraceae (3.3%) and Polygonum (1.6%) are represented uniformly. The major aquatic taxa like, Nymphoides (2.5%), Nymphaea (2.2%), Potamogeton and Typha (2.0% each) are recorded consistently at moderate value. Dipterocarpaceae is represented as major tree constituent (4.2%) followed by Salmalia malabaricum (2.5%), Emblica officinalis (2.1%), Syzygium and Dillenia (2.0% each), Combretaceae (1.9%), Moraceae and Sapotaceae (1.7% each), Ilex (1.6%), Semecarpus (1.5%), Arecaceae and Lagerstroemia (1.3% each), Adina (1.2%) and Meliaceae (1.0%). Mesua and Schima are sporadic. Melastoma malabathricum is represented as lone shrubby element at the value of 1.8% accompanied by Oleaceae (1.7%), Fabaceae (1.5%) and Strobilenthes (1.0%) respectively in lower value. Pinus (4.5%), Betula (3.0%), Alnus (2.8%), Corylus (2.0%), Ulmus and Ericaceae (1.5% each) and Quercus (1.3%) are recovered moderately.

Pollen zone S–IV (40–0 cm)—Grasses–Xanthium– Brassicaceae–Solanaceae–Potamogeton–Dipterocarpaceae– Albizia–Syzygium–Melastoma assemblage: This pollen zone is dated back to $4,270 \pm 70$ years BP at the depth of 40 cm. This part of lithocolumn has 10 cm clay at the base overlain by 30 cm of silty clay with rootlets. This pollen zone is characterized by the predominance of nonarboreals (64.2%) over arboreals (23.1%) and the high land taxa (12.7%). Among nonarboreals, grasses (wild type and cerealia type) are domi-



LEGENDS 0–30-Silty clay with rootlets; 30–40-Clay; 40–70-Silty clay; 70–100-Silty clay with little sand; 100–120-Silty clay with little sand and pebble

Fig. 5-Vegetation composition and climate implication of Lower Subansiri River Basin, Assam.

nant (11.3% and 14.5%) associated with Tubuliflorae (5.7%), Liguliflorae (3.6%), Artemisia (2.8%), Malvaceae (2.4%), Euphorbiaceae and Brassicaceae (1.5% each), Caryophyllaceae (1.4%), Lamiaceae, Convolvulaceae, Acanthaceae and Chenopodiaceae/Amaranthaceae (1.3% each) and Apiaceae (1.0%) are recovered moderately. Impatiens is sporadic. Marshy taxa like Cyperaceae (5.7%) and Onagraceae (1.8%) are recorded moderately. *Polygonum* is registered in sporadic frequency. Aquatics like Nymphaea (1.2%) and Nymphoides (1.1%) are encountered at lower value. However, Potamogeton and Typha are sporadic. Dipterocarpaceae is represented as major constituent (3.0%) followed by Syzygium (1.3%), Combretaceae and Ilex (1.2% each), Salmalia malabaricum, Dillenia and Sapotaceae (1.0% each). Other associated taxa like Emblica officinalis, Arecaceae, Adina, Lagerstroemia, Moraceae, Schima and Mesua ferrea are sporadic. Melastoma *malabathricum* is represented as lone shrubby element (4.0%)accompanied by Oleaceae (1.4%) and Fabaceae (1.1%) in low values. The high land taxa like Pinus (3.1%), Betula (2.5%), Alnus (2.4%), Quercus, Ulmus and Corylus (1.3% each) are recovered consistently at moderate values. Ericaceae is sporadic.

DISCUSSION

The changing frequencies of various arboreal and nonarboreal palynomorph assemblage within soil profile with absolute radiocarbon dates collectively reflect the succession of forest vegetation and corresponding climates within fluctuating monsoonal activity in the study area over a long span of time. Changing assemblage of the aquatic plants at different levels in the pollen sequence significantly infers the lake–level fluctuations/hydrological status of lakes/swamps, which is very much affected by the rainfall regime. Furthermore, the appearance as well as marked alterations in the cerealia pollen along with cultural pollen taxa is taken for the precise understanding of incipient agricultural practice and its subsequent pace in a definite time frame.

During the first phase (12,500–7,960 years BP), the tree–savannah type of vegetation grew in the region under cold and dry climate as is evident by the higher value of terrestrial nonarboreal elements, viz. grasses, Tubuliflorae, Liguliflorae and Apiaceae along with scattered open mixed deciduous and semievergreen elements like Combretaceae, Lauraceae, *Salmalia malabaricum*, *Lagerstroemia* and *Ilex* together with stray presence of Dipterocarpaceae. The occurrence of *Potamogeton*, *Nymphaea* and *Nymphoides* suggests the existence of perennial waterlogged condition during this phase. The high land taxa comprised of conifers (*Pinus*) and broad leaved elements (*Alnus*, *Corylus* and *Ulmus*) are growing in higher reaches of Arunachal Himalaya proves wind activity. The stable presence of the taxa, viz. *Tsuga* and Ericaceae suggests water flowing conditions from high altitude during this period.

Subsequently during the second phase (7,960–6,421 years BP), the mixed deciduous forest got enriched with the



Fig. 6-Estimation of monsoonal activity fluctuation by using pollen marker taxa.

immigration of some more tropical deciduous arboreals like Semecarpus anacardium, Emblica officinalis and Dillenia pentagyna. Dipterocarpaceae got colonized as is evident by a relative rise in its frequency. Sharp declining trends in frequency of most of the terrestrial nonarboreal taxa have been observed. However, both marshy and aquatics especially Onagraceae, Impatiens and Potamogeton increased their value than the preceding phase. This vegetation change clearly depicts the formation of deciduous mixed forest along with vast stretches of swampy areas with wider perennial water logged conditions. The high land taxa (conifers and broad leave taxa) show continuously stable occurrence in the palynoassemblage in the sediment. The overall pollen assemblage suggests the onset of warm and humid climatic condition during this period. Pastoral activity is marked by the presence in traces of cerealia and culture pollen taxa in the palynoassemblage.

During the third phase (6,421–4,270 years BP), the mixed deciduous forest got more enriched with dense tropical mixed deciduous arboreals as compared to the preceding phase. The arboreal taxa are same as the preceding phase but with their higher density, viz. Dipterocarpaceae, *Syzygium cumunii*, Arecaceae and Moraceae. The semievergreen taxa such as *Mesua ferrea* and Combretaceae are consistent as the preceding phase. The high land taxa are registered in persistent frequency. The frequencies of aquatic taxa are increased than the preceding phase and supporting good southeast monsoon. The overall pollen assemblage is suggestive of the warm and humid climatic condition during this phase.

During the fourth phase (4,270 years BP onwards), mostly the semievergreen and deciduous arboreals such as Mesua ferrea, Dipterocarpaceae, Arecaceae and Syzygium cumunii have decreased whereas, the terrestrial nonarboreal taxa have increased as compared to those present during the preceding phase. The decreased values of most of the aquatic pollens indicate decreased water logged conditions as well as weakened southeast monsoon than the preceding phase. Sudden decline in Onagraceae and Impatiens are also indicator of low precipitation. The shrubby taxa Melastoma malabathricum shows its peak during this phase suggesting the declining monsoonal activity and dryness during the phase. The sudden increase of Cerealia, Brassicaceae, Chenopodiaceae, Apiaceae and Xanthium implies high pastoral activity by the new settlers in and around the study area. The consistent presence of high land taxa supports long distance pollen transport along with high wind activity. The overall scenario is suggestive of the deteriorated mixed deciduous forest under warm and relatively dry climatic condition during this phase (Fig. 5).

Based on the investigations in lakes Didwana, Sambhar and Lunkaransar (Singh *et al.*, 1974, 1990), identified the following ecologically sensitive pollens, (i) *Artemisia*, an element of the > 50 cm annual rainfall regime, (ii) *Oldenlandia*, which is indicative of higher precipitation as it is found in sub humid and humid regions of NW India, (iii) *Typha*, a fresh water swamp taxon and (iv) *Syzygium cumunii*, indicative of > 850 mm average annual rainfall.

In this paper we have considered the marker pollen taxa to estimate monsoonal activity. The plant taxa Dipterocarpaceae, Syzygium cumunii, Arecaceae, Moraceae, Impatiens and aquatic taxa (Nymphoides, Nymphaea, Potamogeton and Typha) are considered as the high monsoonal marker pollen taxa and Melastoma malabathricum as the low monsoonal activity marker taxa to estimate the monsoonal activity (Singh et al., 1990; Barboni & Bonnefille, 2001), (Fig. 6). Results indicate that there was a fluctuation of monsoonal activity during the recent past. During the period 12,500 to 7,960 years BP, the monsoonal activity was relatively low as evidenced by low frequency of the high monsoonal activity marker taxa (Fig. 4). However, between 7,960 to 6,421 years BP, the monsoonal activity is increased than the preceding period as evidenced by increased frequency of marker pollen taxa in the pollen assemblage. From 6,421 to 4,270 years BP, the monsoonal activity is considerably higher than the preceding phase as marked by the high frequency of marker taxa, viz. Dipterocarpaceae, Syzygium cumunii, Arecaceae and Moraceae. Accordingly, during 4,270 years BP to present, the monsoonal activity relatively declined as suggested by the dominance of Melastoma malabathricum (sharply increased). The other marker taxa, viz. Dipterocarpaceae, Syzygium cumunii, Arecaceae and Impatiens also show lower percentage in the palynoassemblage thereby strongly supporting the weakened monsoonal activity.

It is quite important to note that since around 2,000 years BP, the monsoonal activity is strengthened in Indian Subcontinent, Arabian Sea and western Tibet (Phadtare, 2000). The carbon and oxygen isotopic records from the stalagmite in Xiangshui Cave, Guilin suggest that the Asian Monsoon was relatively stronger from 6,000 to 3,800 years BP and gradually weakening during the interval. However, since 3,800 years BP onwards, there is a relatively weak Asian Monsoon (Zhang *et al.*, 2004). The South Asian monsoon in Holocene records reveal substantial regional differences (Staubwasser, 2006). More investigation with multiproxy palaeoclimate data are needed to find the precise and regional understanding of the palaeoclimate and vegetation in and around the study area.

Acknowledgements—We thank the Director, Birbal Sahni Institute of Palaeobotany (BSIP), Lucknow for infrastructure facility and permission to publish the paper. We also thank Mrs Indra Goel, Technical officer, BSIP for chemical analysis of the samples. Thanks are also due to forest officials of Dhemaji District, Assam for providing the necessary facilities during collection of samples.

REFERENCES

- Barboni D & Bonnefille R 2001. Precipitation signal in modern pollen rain from tropical forests of south India. Review of Palaeobotany & Palynology 114: 239–258.
- Basumatary SK & Bera SK 2007. Modern pollen–spore assemblage from sediment of tropical moist deciduous forest, East Garo Hills, Meghalaya. Journal of Palynology 43: 111–118.

- Basumatary SK & Bera SK 2010. Development of vegetation and climate change in West Garo Hills since late Holocene: pollen sequence and anthropogenic impact. Journal of Indian Botanical Society 89: 143–148.
- Bera SK 2000. Modern pollen deposition in Mikir Hills, Assam, India. Palaeobotanist 49: 325–328.
- Bera SK 2003. Early Holocene pollen data from Mikir Hills, Assam, India. Palaeobotanist 52: 121–126.
- Bera SK, Dixit S, Basumatary SK & Gogoi R 2008. Evidence of biological degradation in sediments of Deepor Beel Ramsar site, Assam as inferred by degraded palynomorphs and fungal remains. Current Science 95: 178–181.
- Bera SK, Basumatary SK, Nautiyal CM, Dixit S, Mao AA & Gogoi R 2010. Late Holocene climate and vegetation change in Dzuko Valley, north east India. Journal of the Palaeontological Society of India 56: 143–148.
- Bhattacharya K & Chanda S 1986. Quaternary pollen analysis of a peat sample from Gangtok (Sikkim). Science Culture 52: 139–140.
- Bhattacharya K & Chanda S 1988. Late Quaternary vegetational history, palaeoecology and biostratigraphy of some deposits of Brahmaputra Basin, Upper Assam, India. Journal of Palynology 23–24: 225–237.
- Bhattacharya K & Chanda S 1992. Late–Quaternary vegetational history of Upper Assam, India. Review of Palaeobotany & Palynology 72: 325–333.
- Bhattacharyya A, Sharma J, Shah SK & Chaudhury V 2007. Climatic changes last 1800 years BP from Paradise Lake, Sela Pass, Arunachal Pradesh, northeast Himalaya. Current Science 93: 983–987.
- Champion HG & Seth SK 1968. A revised Survey of the Forest Type of India, Delhi.
- Chauhan MS & Mandaokar BD 2006. Pollen proxy records of vegetation and climate change during recent past in southern Mizoram, India. Gondwana Geological Magazine 21: 115–119.
- Chauhan MS & Sharma C 1996. Late Holocene vegetation of Darjeeling (Jore–Pokhari), eastern Himalaya. Palaeobotanist 45: 125–129.
- D' Costa M & Mukharjee BB 1986. Holocene history of ferns in Darjeeling Hills, eastern Himalaya. Phytomorphology 36: 151–163.
- Demske D, Tarasov PE, Wunnemann B & Riedel F 2009. Late glacial and Holocene vegetation, Indian Monsoon and westerly circulation in the Trans–Himalaya recorded in the lacustrine pollen sequence from Tso Kar, Ladakh, NW India. Palaeogeography, Palaeoclimatology, Palaeoecology 279: 172–185.
- Ely LL, Enzel Y, Baker VR & Cayan DR 1993. A 5000 year record of extreme floods and climate change in southwestern United State. Science 262: 410–412.
- Erdtman G 1953. An Introduction to Pollen Analysis, Waltham, Mass; USA. Faegri K & Iversen J 1964. Text book of Pollen Analysis (IV ed.), Wiley.
- Fleitmann D, Burns SJ, Mudelsee M, Neff U, Kramers J, Mangini A & Matter A 2003. Holocene forcing of the Indian Monsoon recorded in a stalagmite from southern Oman, Science 300: 1737–1739.
- Goodbred SL Jr. & Kuehl SA 2000. Enormous Ganges–Brahmaputra sediment discharge during strengthened early Holocene Monsoon. Geology 28: 1083–1086.
- Goswami AB 1981. Palynological and radiocarbon dating of peat deposits in Tripura. *In:* Khosla SC & Kacchara RP (Editors)—Proceeding of 9th Indian Colloquium Micropalaeontology Stratigraphy, Udaipur: 192–200.
- Goswami DC 1997. Brahmaputra River Basin forestry. In: forestry and key Asian watersheds, Myint AK, Hofer T (Editors). ICIMOD: Katmandu: 32–37.
- Goswami U, Sarma JN & Patgiri AD 1990. River channel changes of the Subansiri in Assam, India. Geomorphology 30: 227–244.
- Gupta HP 1971. Pollen analysis investigations of some Pleistocene samples from Tocklai, Cinnamara, Assam. Palaeobotanist 18: 234–236.
- Gupta HP & Sharma C 1985. Pollen analysis of modern sediments from Khasi and Jaintia Hills, Meghalaya, India. Journal of Palynology 21: 167–173.
- Gupta AK, Anderson DM & Overpeck JT 2003. Abrupt changes in the Asian southwest monsoon during the Holocene and their links to the North Atlantic Ocean. Nature 421: 354–357.

- Gupta AK, Anderson DM, Pandey DN & Singhvi AK 2006. Adaptation and human migration and evidence of agriculture coincident with changes in the Indian Summer Monsoon the Holocene. Current Science 90: 1082–1090.
- Karunakaran C 1974. Geology and mineral resources of the states of India. Part IV. Arunachal Pradesh, Assam, Manipur, Meghalaya, Nagaland and Tripura. Geological Society of India, Miscellaneous Publication 30: 1–124.
- Kumaran KPN, Nair KM, Shindikar M, Limaye RB & Padmalal D 2005. Stratigraphical and palynological appraisal of the late Quaternary mangrove deposits of the west coast of India. Quaternary Research 64: 418–431.
- Mandaokar BD, Chauhan MS & Chatterjee S 2008. Fungal remains from Late Holocene Lake deposit of Demagiri, Mizoram, India and their palaeoclimatic implications. Journal of Palaeontological Society of India 52: 197–205.
- Meher–Homji VM & Gupta HP 1999. A critical appraisal of vegetation and climate changes during Quaternary in the Indian region. PINSA B65: 205–244.
- Mehrotra RC, Liu, Xiu–Qun Li, Cheng–Sen, Wang Yu–Fei & Chauhan MS 2005. Comparison of the Tertiary flora of southwest China and northeast India and its significance in the antiquity of the modern Himalayan flora. Review of Palaeobotany & Palynology 135: 145–163.
- Mehrotra RC, Bera SK, Basumatary SK and Srivastava G 2011. Study of fossil wood from the Middle–Late Miocene sediments of Dhemaji and Lakhimpur districts of Assam, India and its palaeoecological and palaeophytogeographical implications. Journal of earth system science 120: 681–702.
- Nautiyal CM & Chauhan MS 2009. Late Holocene vegetation and climate change in Loktak Lake region, Manipur, based on pollen and chemical evidence. Palaeobotanist 58: 21–28.
- Phadtare NR 2000. Sharp decrease in summer monsoon strength 4000–3500 cal yr B. P. in the central higher Himalaya of India based on pollen evidence from Alpine peat. Quaternary Research 53: 122–129.
- Prasad MNV & Ramesh NR 1983. Fungal remains from the Holocene sediments of Tripura, India. Current Science 52: 254–256.
- Prasad MNV 1986. Fungal remains from the Holocene peat deposits of Tripura State, north eastern India. Pollen spores 28: 365–390.
- Prasad MNV 1988. Ecological and archaeological aspects of Holocene deposits of Tripura State, north eastern India. Proceeding of Indian National Science Academy 54: 452–460.
- Prasad S & Enzel Y 2006. Holocene palaeoclimate of India. Quaternary Research 66: 442–453.
- Roy P & Chanda S 1987. Late Quaternary vegetational history and biostratigraphy of Loktak Lake of Manipur, India. Transactions of Bose Research Institute 50: 73–80.
- Sharma C & Chauhan MS 2001. Late Holocene vegetation and climate of Kupur (Sikkim) Eastern Himalaya, India. Journal of Palaeontological Society of India 46: 51–58.
- Singh G, Joshi RD, Chopra SK & Singh AB 1974. Late Quaternary history of vegetation and climate of the Rajasthan Desert, India. Philosophical Transactions of the Royal Society, London (Biological Science) 267: 467–501.
- Singh G, Wasson RJ & Agarwal DP 1990. Vegetational and seasonal climatic changes since the last full glacial in the Thar Desert, northwestern India. Review of Palaeobotany & Palynology 64: 351–358.
- Staubwasser M 2006. An overview of Holocene South Asian Monsoon records–Monsoon domains and regional contrasts. Journal of Geological Society of India 68: 433–446.
- Swain AM, Kutzbach JE & Hastenraths S 1983. Estimates of Holocene precipitation for Rajasthan, India, based on pollen and Lake–Level data. Quaternary Research 19: 1–17.
- Zhang M, Yuan D, Lin Y, Qin J, Bin L, Cheng H & Edwards L 2004. A 6000–year high–resolution climatic record from a stalagmite in Xiangshui Cave, Guilin, China. The Holocene 14: 697–702.